

# Attenuation of high-frequency body waves and its anisotropy in the External Dinarides

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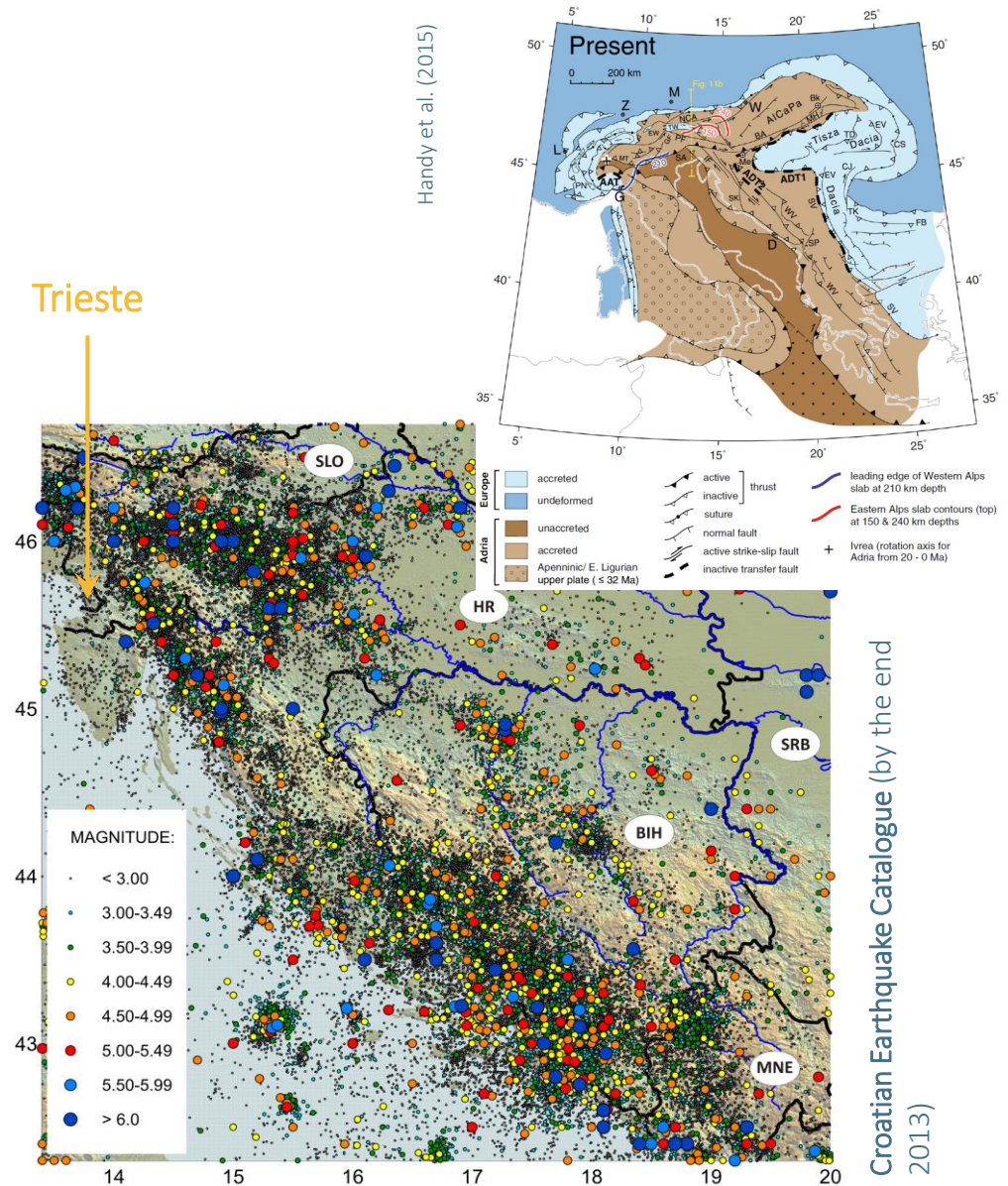
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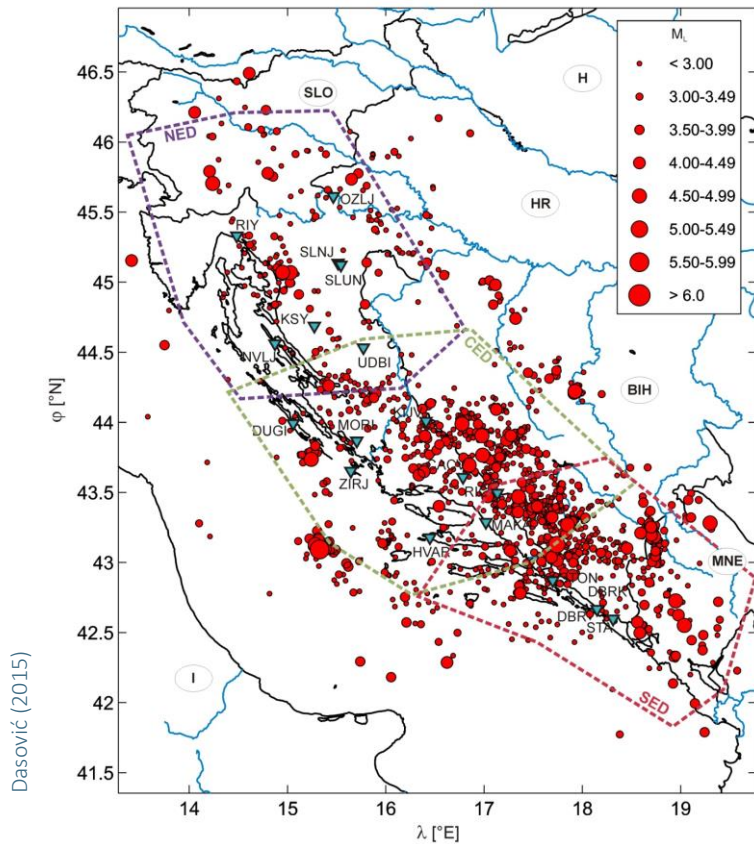
# Why?

- The (External) Dinarides → complex, seismically active and *under-researched*
- High-frequency (1–24 Hz) body waves attenuation → local earthquake → structure of the crust, moment magnitude, seismic hazard
- Methods:
  - (Extended) coda normalization method →  $Q_p$  and  $Q_s$
  - Multiple scattering: MLTWA method →  $Q_i$  and  $Q_{sc}$



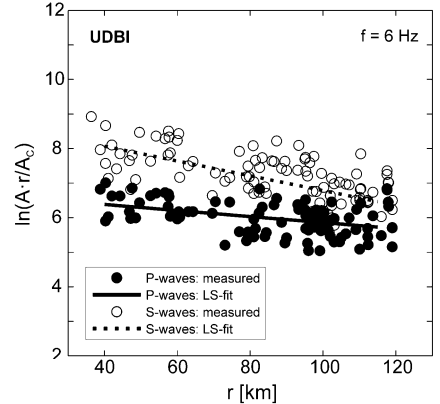
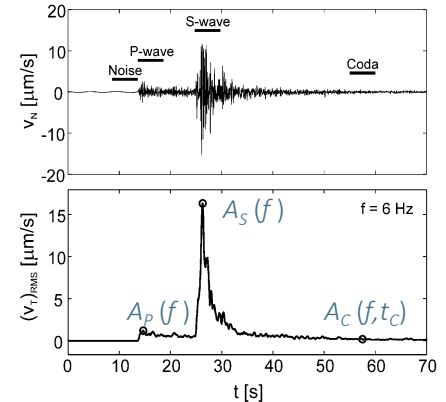
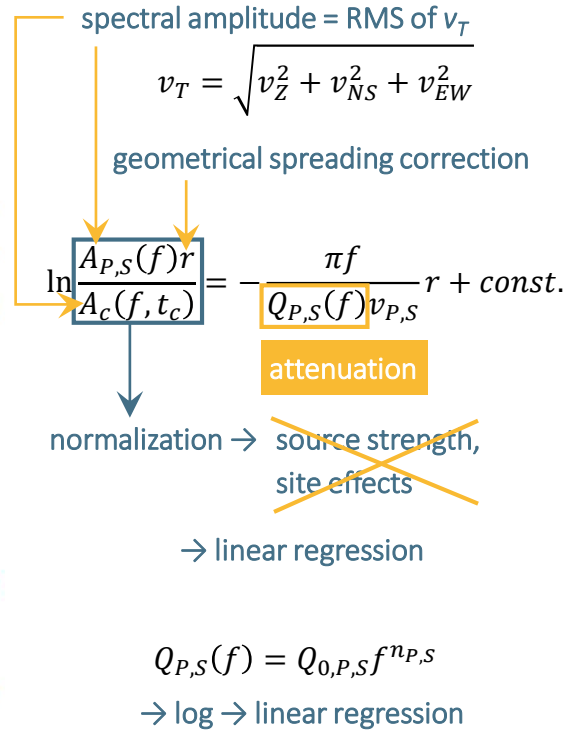
# (Extended) Coda normalization method

- Aki (1980) + Yoshimoto et al. (1993)
- $Q_{P,S}$  → direct P- and S-waves → (upper) crust

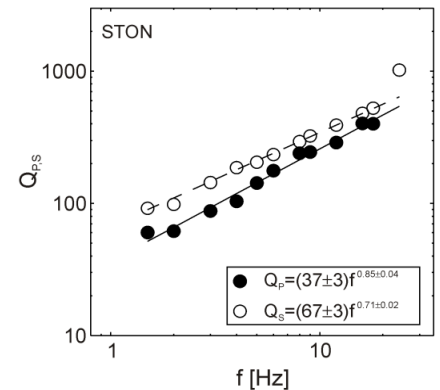


Dasović (2015)

17 BB stations,  $40 \leq D \leq 120$  km,  $M_L \geq 2.0$ , 2002–2014:  
1526 EQ, 2558 seismograms

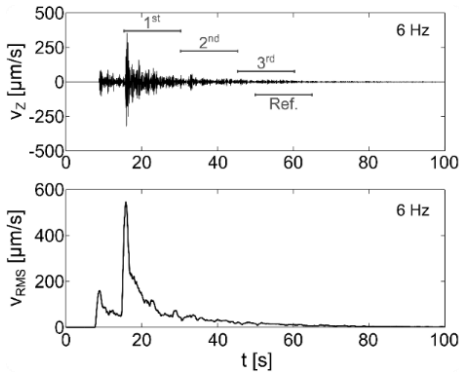


Dasović et al. (2015)



# MLTWA method

- Hoshiba et al. (1991), Fehler et al. (1992), Hoshiba (1993)
- $Q_{sc} + Q_i \rightarrow$  separate estimates!



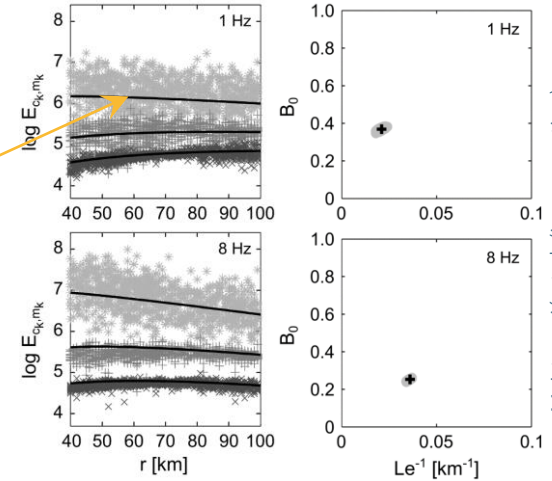
Majstorović et al. (in preparation)

theoretical energy  $\rightarrow f(B_0, Le^{-1})$   
 (isotropic multiple scattering homogeneous  
 halfspace + radiative transfer theory)

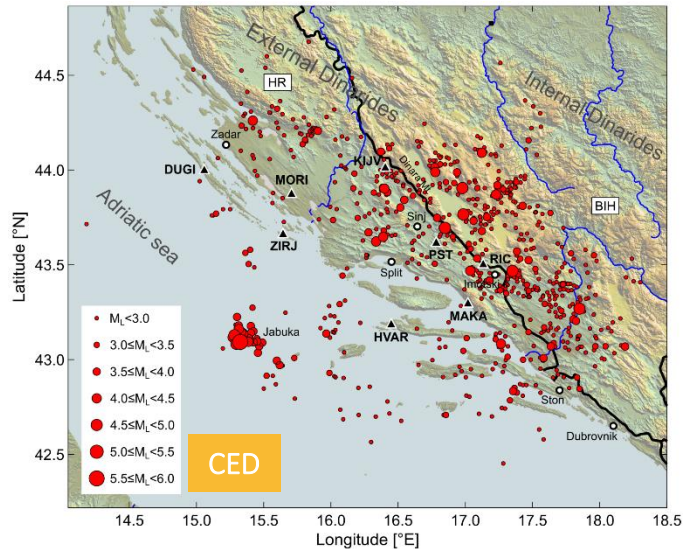
vs.

observed energy

- \* 1<sup>st</sup> window: 0–15 s
- + 2<sup>nd</sup> window: 15–30 s
- × 3<sup>rd</sup> window: 30–45 s



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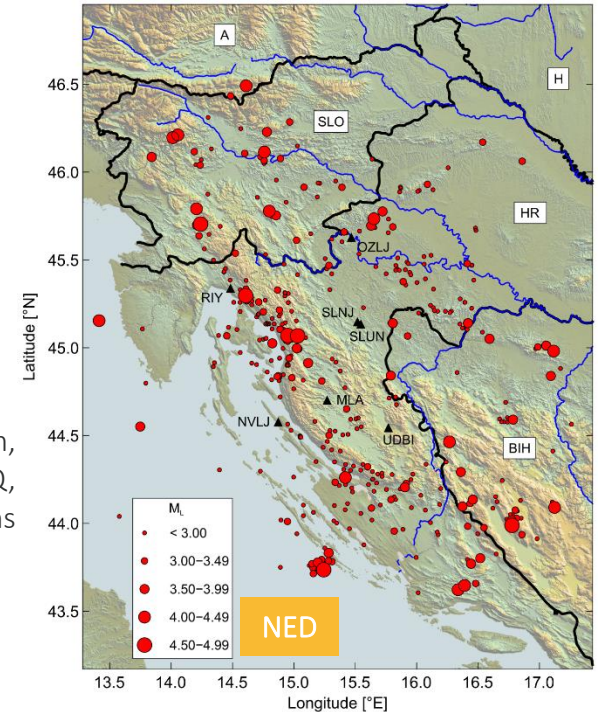
$$B_0 = \frac{\eta_{sc}}{\eta_{sc} + \eta_i} \quad \text{seismic albedo}$$

$$Le^{-1} = \eta_{sc} + \eta_i \quad \text{total attenuation}$$

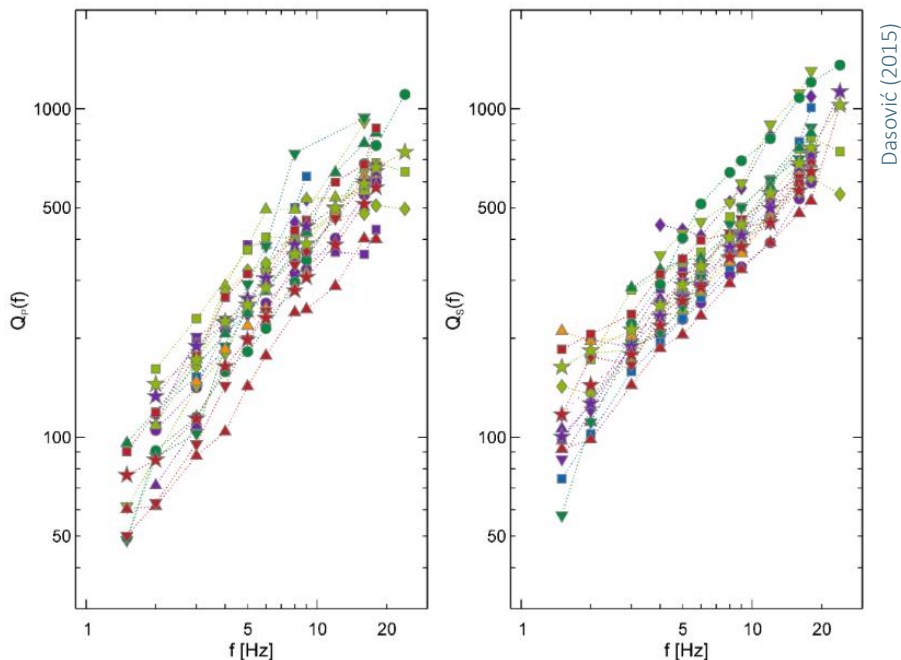
$$\eta_{i,sc} = \frac{2\pi f c}{v_s Q_{i,sc}}$$

6 BB stations,  $10 \leq D \leq 120$  km,  
 $ML \geq 2.0$ , 2003–2015: 415 EQ,  
 717 seismograms

8 BB stations,  $40 \leq D \leq 120$  km,  
 $ML \geq 2.0$ , 2002–2014: 750 EQ,  
 985 seismograms



# (Extended) Coda normalization method

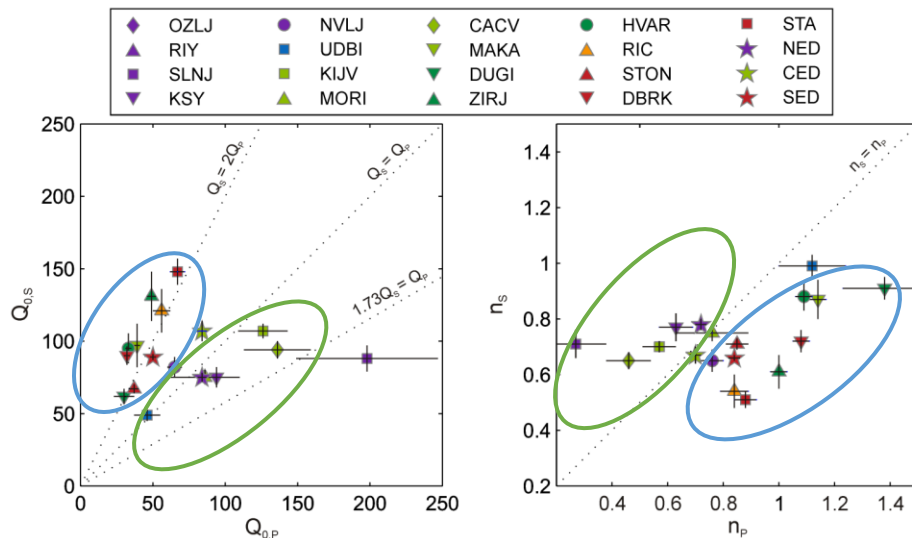


Dasović (2015)

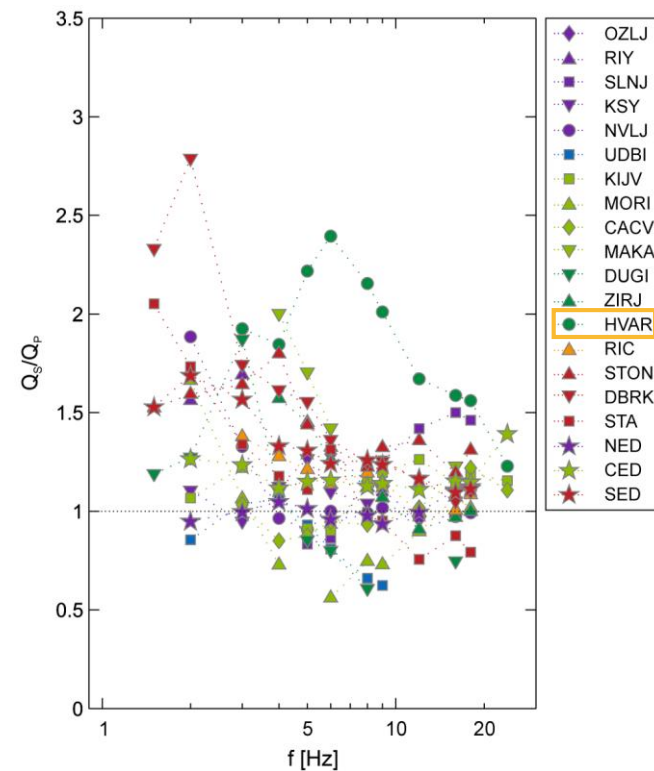
$Q_s / Q_p \rightarrow$  scattering, fluid saturation

$Q_s / Q_p \geq 1 \rightarrow$  strong scattering, partially saturated rocks

$Q_s / Q_p < 1 \rightarrow$  weaker scattering, fully saturated or dry rocks



Dasović (2015)

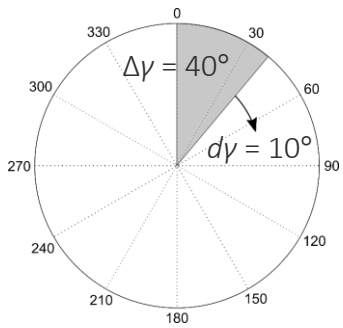


islands CED + SED

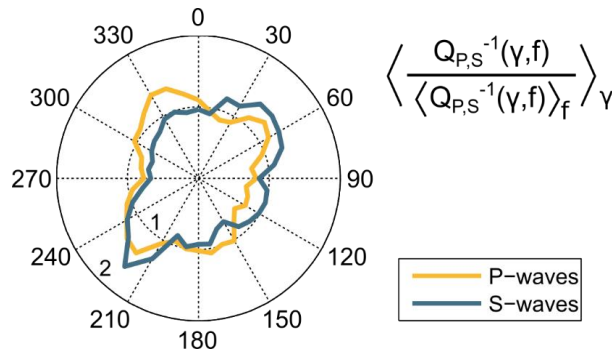
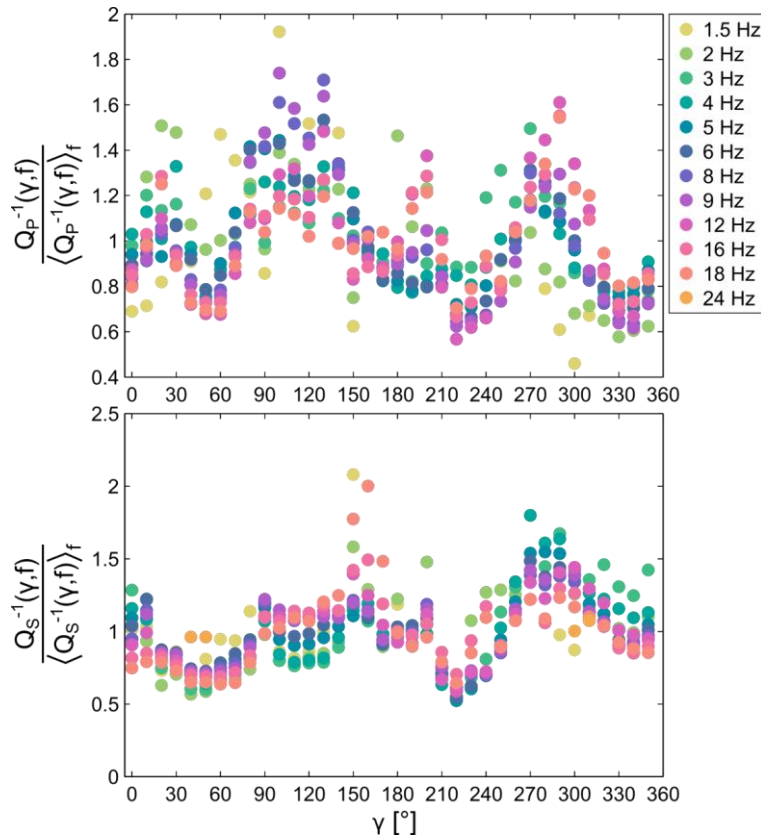
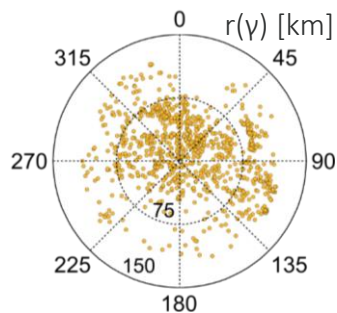
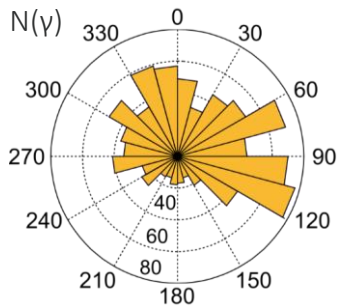
continent CED + NED

Dasović (2015)

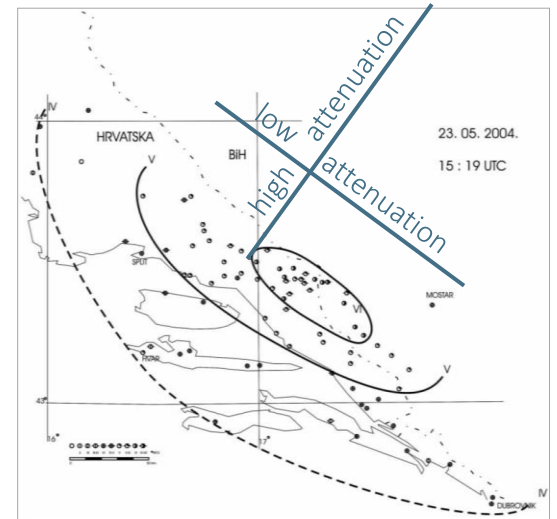
# (Extended) Coda normalization method



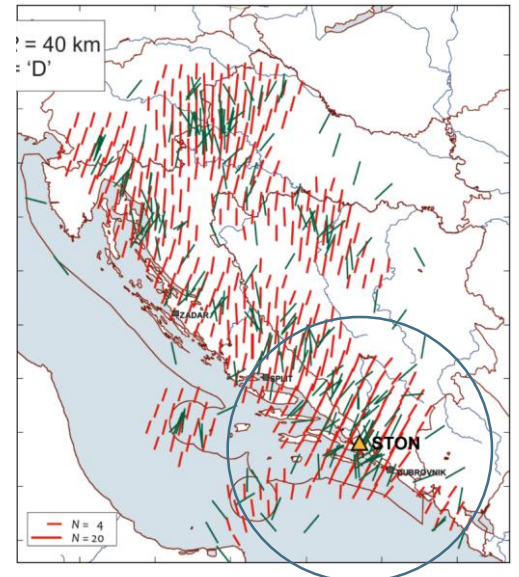
STON:  $D \leq 120$  km,  
 $ML \geq 2.0$ , 2003–2015:  
 940 EQ



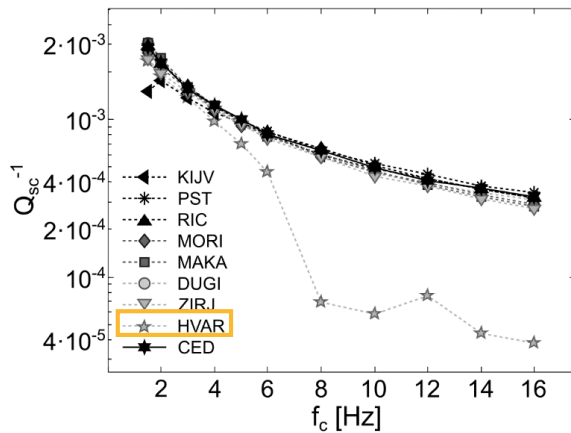
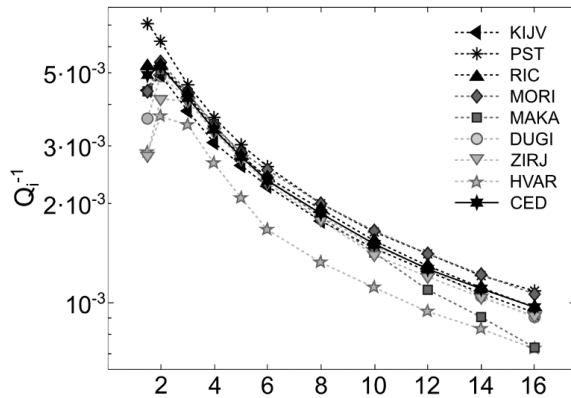
$$\left\langle \frac{Q_{P,S}^{-1}(\gamma, f)}{\langle Q_{P,S}^{-1}(\gamma, f) \rangle_f} \right\rangle_\gamma$$



→ Attenuation anisotropy!



# MLTWA method

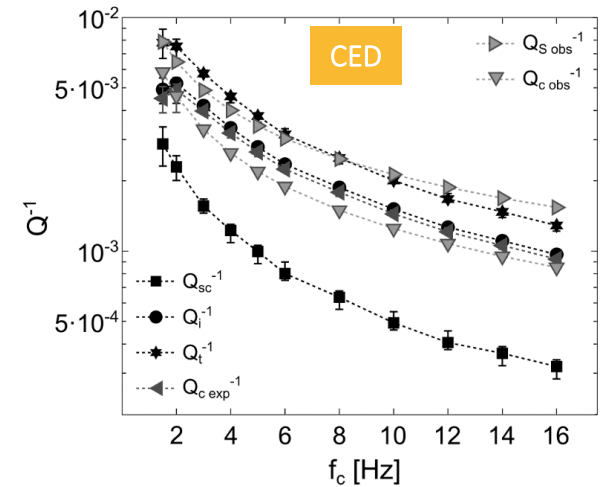


$$Q_i^{-1} > Q_{sc}^{-1}$$

$$Q_t^{-1} \approx Q_{S,obs}^{-1}$$

$$Q_{C,exp}^{-1} \approx Q_{C,obs}^{-1}$$

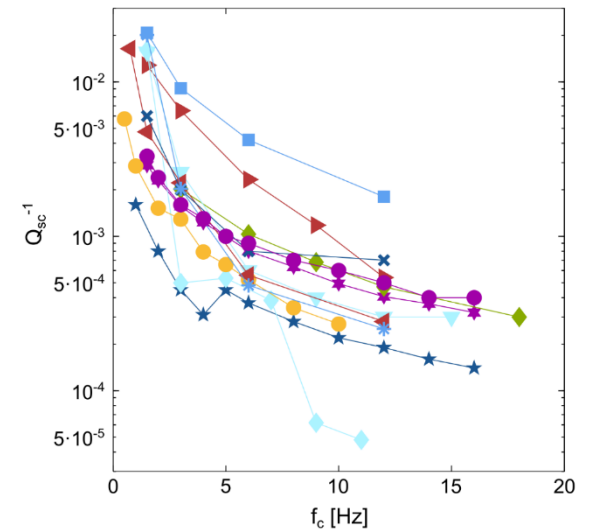
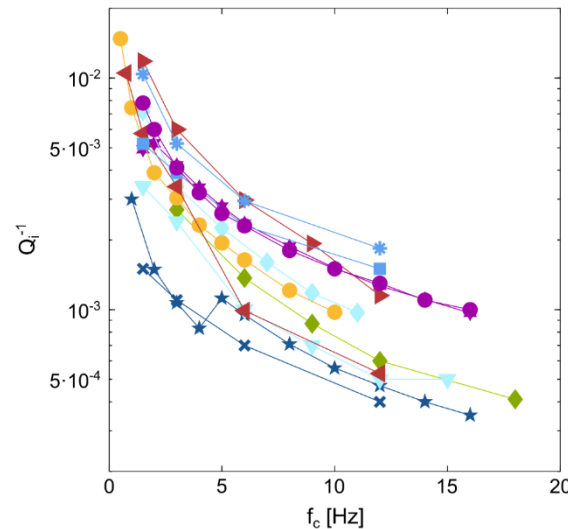
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- ◆ CED
- NED
- ★ Friuli (NE Italy)
- ★ NC Italy
- ★ Umbria-Marche (C Italy)
- Gran Sasso Mt. (C Italy)
- ◆ S Apennines (Italy)
- ◆ SE Sicily (Italy)
- ▲ SW Anatolia (Turkey)
- ▶ E Turkey
- Israel
- ◆ Cairo (Egypt)

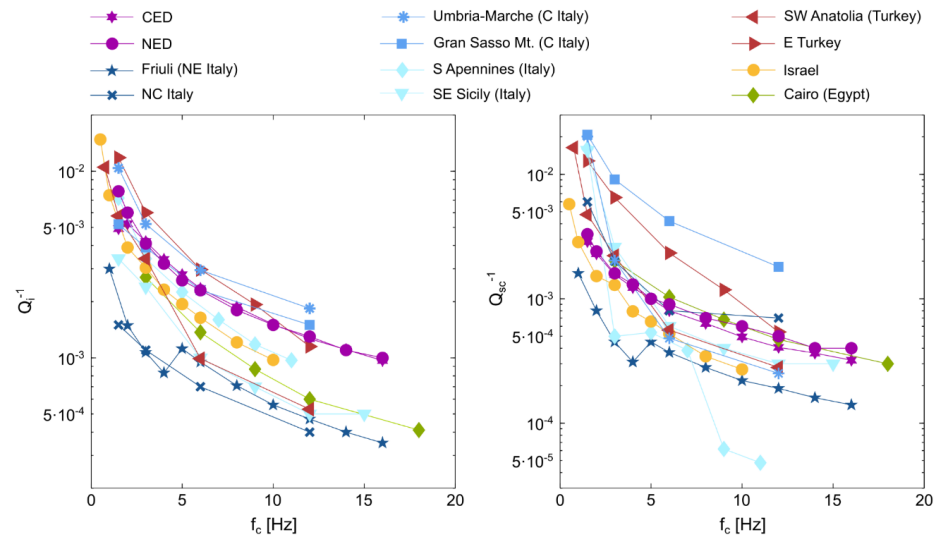
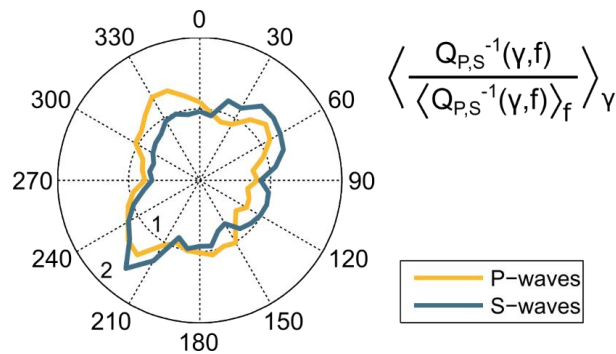
Majstorović et al. (in preparation)

$B_0 < 0.5 \rightarrow$  intrinsic attenuation dominates  
 $L_e = 4.8 \text{ km (1.5 Hz)} - 27.0 \text{ km (16 Hz)} \rightarrow$  CED



## To conclude...

- Attenuation of high-frequency body waves is high in the External Dinarides
- In general: P-wave attenuation > S-wave attenuation ( $Q_p < Q_s$ )
- Intrinsic attenuation dominates over scattering attenuation ( $Q_i^{-1} > Q_{sc}^{-1}$ )
- Anisotropy of the attenuation of P- and S-waves
  - high attenuation parallel to P-axis
  - Low attenuation parallel to the strike of the Dinarides





## References

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## Aknowledgments

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